

Meeting EPIGRAM 2011  
30 may-1 Juin. Ile de Ré

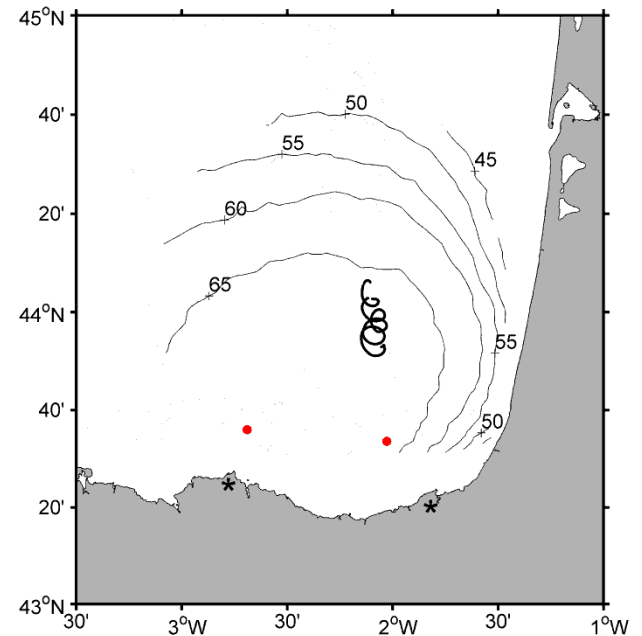
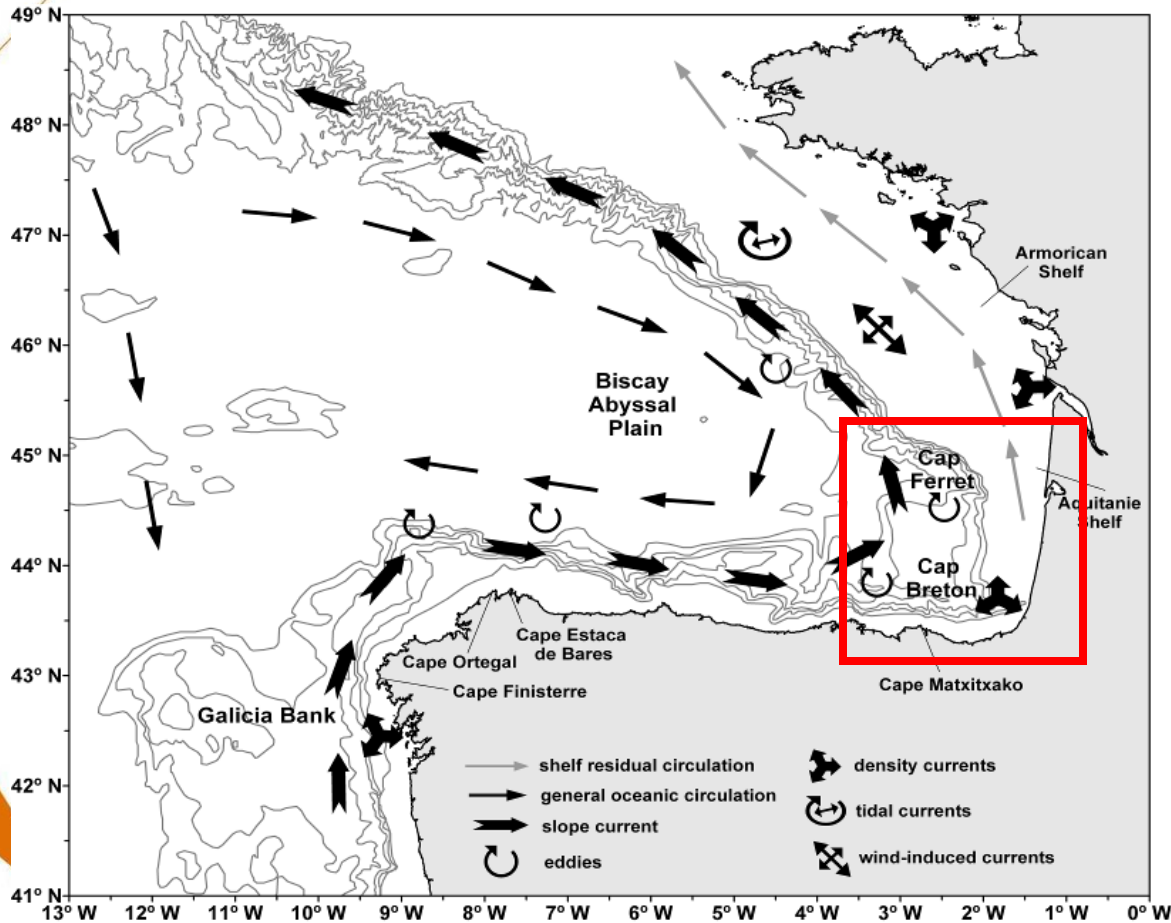
# Mapping near-inertial variability in the SE Bay of Biscay from HF radar data and two offshore moored buoys

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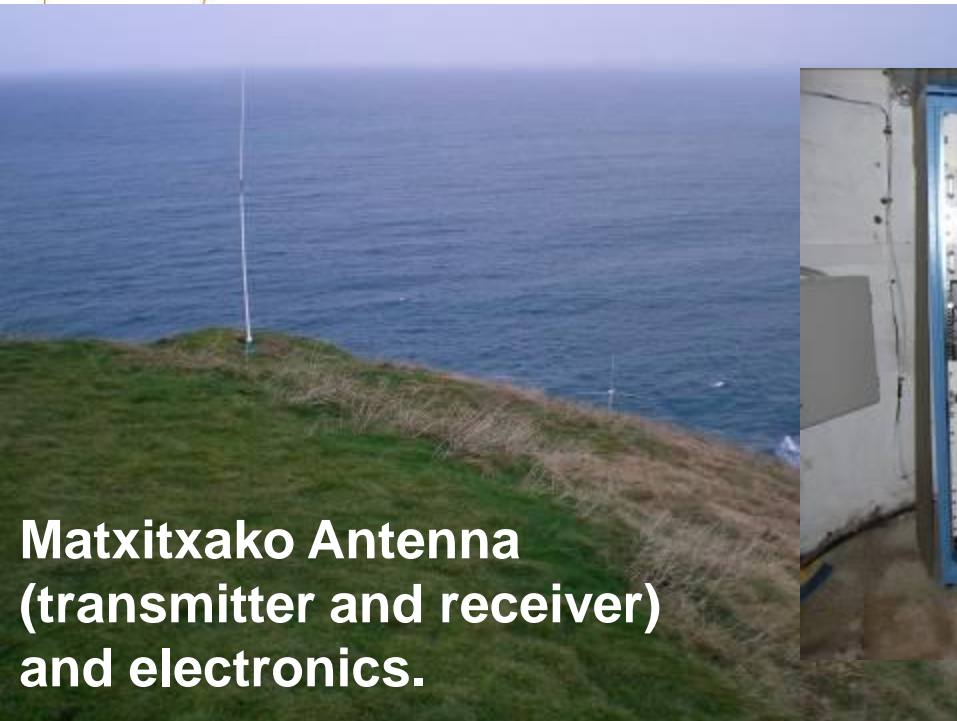
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## Mean circulation patterns



Study area and available data (BOLD ISOLINES: availability of radar data (%) for the target year 2009 )

# Introduction: the HF radar system main characteristics



Two radial site stations: Ilger (Donostia) and Matxitxako  
Central combine site in Vitoria

**WORKING SINCE 2009**

## MAIN CHARACTERISTICS of the BASQUE SYSTEM

Radar Frequency (MHz)	Radar Wavelength (m)	Ocean Wavelength (m)	Depth of Current <sup>1</sup> (m)	Typical Range <sup>2</sup> (km)	Typical Resolution (km)	Typical Bandwidth (kHz)	Upper H <sub>1/3</sub> Limit <sup>4</sup> (m)
4.86	60	30	2-3	~150	~ 5	30	25

# Introduction: the HF radar system main characteristics

## FROM CROSS SPECTRA TO RADIAL VELOCITIES

### **MUSIC (Multiple Signal Classification) Algorithm**

[Schmidt, R.O., (1986), *IEEE Trans. Antennas Propagat.*, vol. AP-34, pp. 276-280]

- Linear solutions based on eigen-analysis of averaged spectral covariances
- Capable of more than two solutions with little added numerical burden
- Chooses the number of bearings based on both eigenvalue & amplitude ratios
- Capable of incorporating measured, distorted antenna patterns

***OUTPUT: Vectors in polar coordinate system centered at receiving antenna:***

- ✓ ***1 radial map per averaged cross spectra file***
- ✓ ***Typically, seven radial maps “merged” into one hourly map (3-hours running average)***
- ✓ ***Angular resolution is 5°***

# Introduction: the HF radar system main characteristics

## FROM RADIAL TO TOTAL VELOCITIES

Several radials from each site contribute to each total vector, being resolved into orthogonal components (U, V)

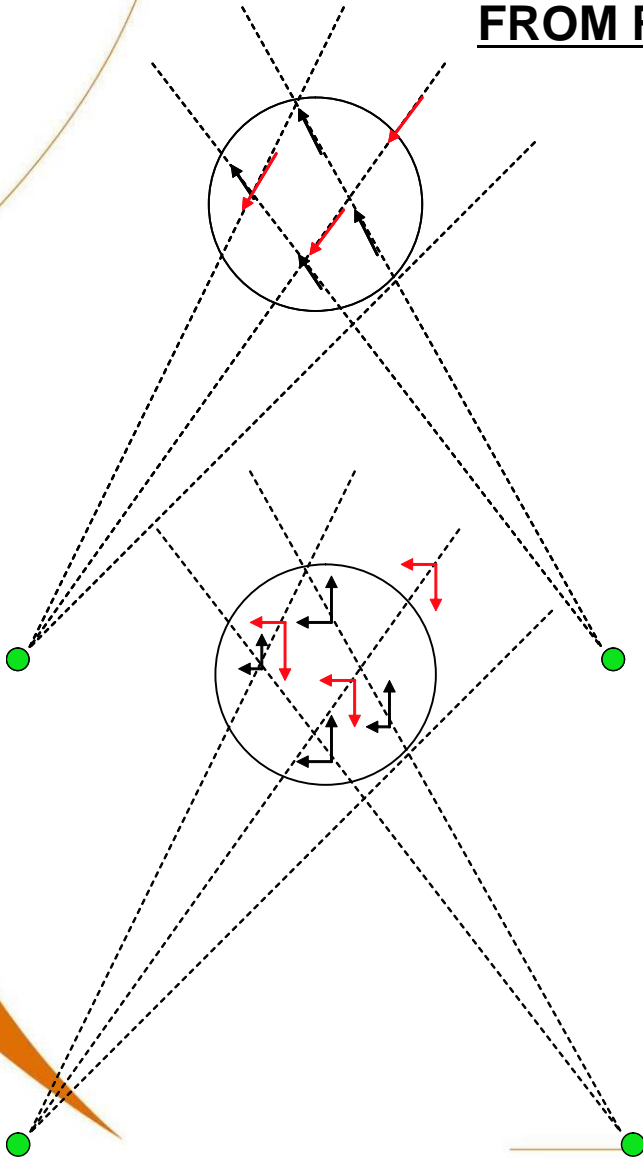
Linear least squares performed on U, V components separately

From radials with:

- An **angle of incidence** greater than  $\alpha$   
 $\alpha \text{ min} = 30^\circ$
- And **total velocities**  $< 120 \text{ cm/s}$

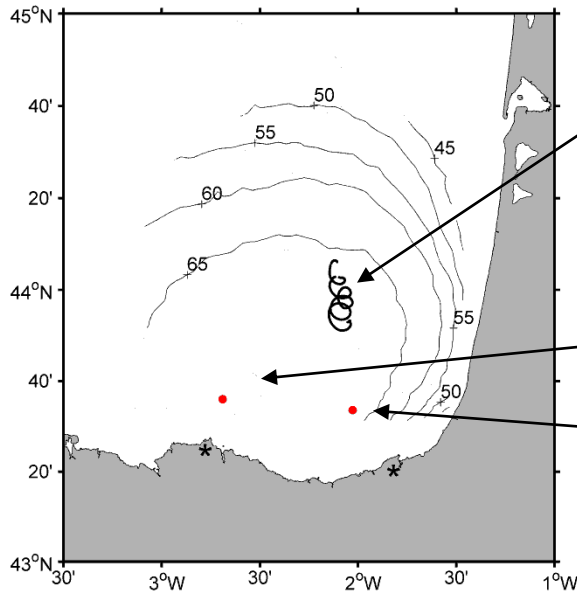
Defining a

- **GRID SPACING:** 5km
- **AVERAGING RADIUS:** of 20 KM



# HF radar vs. In-situ data

**HF RADAR (4.5 MHz) integrates currents vertically within the first 2-3 m of the water column. Radial resolution= 5km; angular resolution 5°; 3h running window hourly data.**

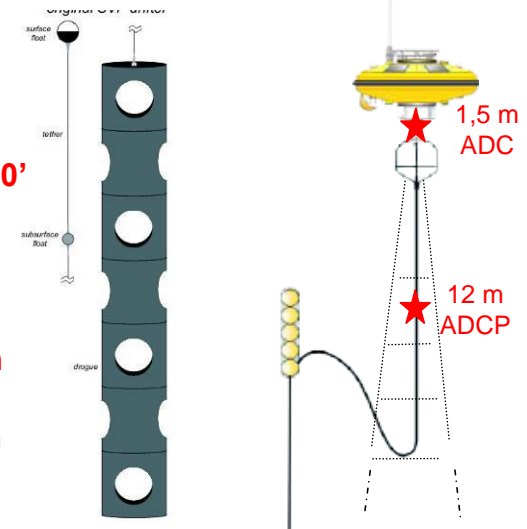


Drifter -

Centred at 8 m  
10 m long drogue  
GPS positions every 30'  
3h running mean

Matxitxako  
Donostia

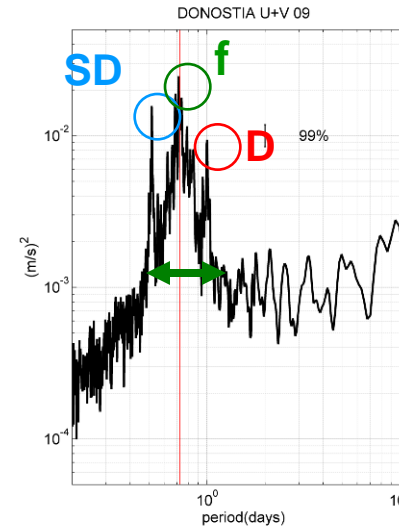
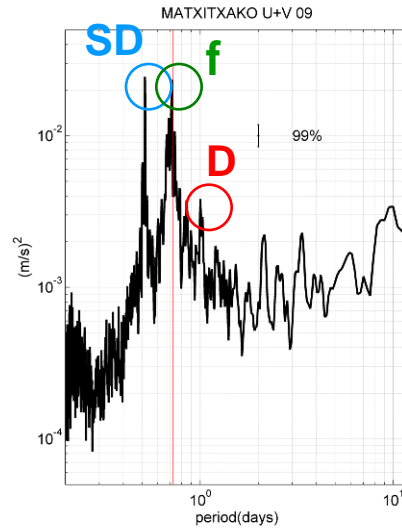
At 1.5 –and 12 m  
Hourly data,  
3h running mean



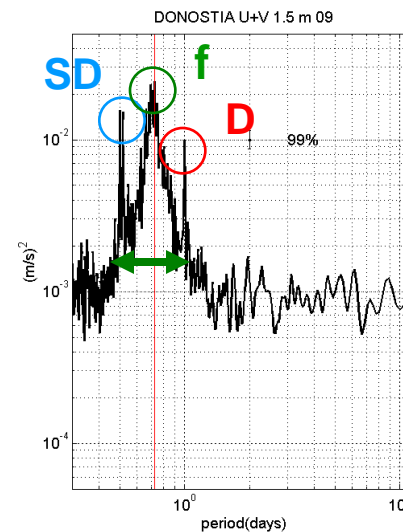
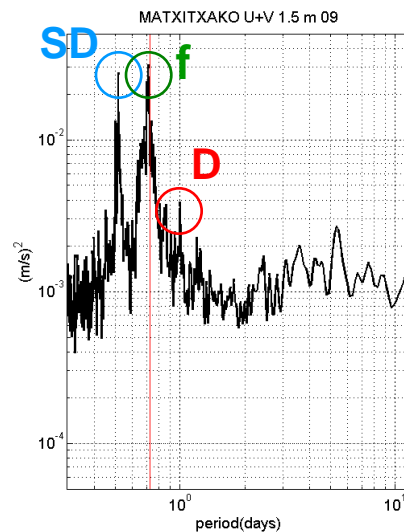
	R u, v	RMS [ $\text{cm}\cdot\text{s}^{-1}$ ] u, v	MRD u, v
Matxitxako 1.5 m vs. radar	0.86, 0.64	8.10, 8.12	-1.10, -0.92
Matxitxako 12 m vs. radar	0.74, 0.50	10.79, 9.11	-1.85, -2.70
Donostia 1.5 m vs. radar	0.53, 0.35	10.38, 12.89	-1.31, 0.44
Donostia 12 m vs. radar	0.37, 0.21	10.31, 13.40	0.84, -0.22
Drifter vs. radar	0.72, 0.91	16.54, 9.22	-0.22, -0.01
Matxitxako 1.5 m vs. 12 m	0.80, 0.60	9.96, 8.92	1.97, 0.60
Donostia 1.5 m vs. 12 m	0.37, 0.28	11.91, 11.61	-1.36, -0.83

Idea of vertical shear

# HF radar vs. In-situ data



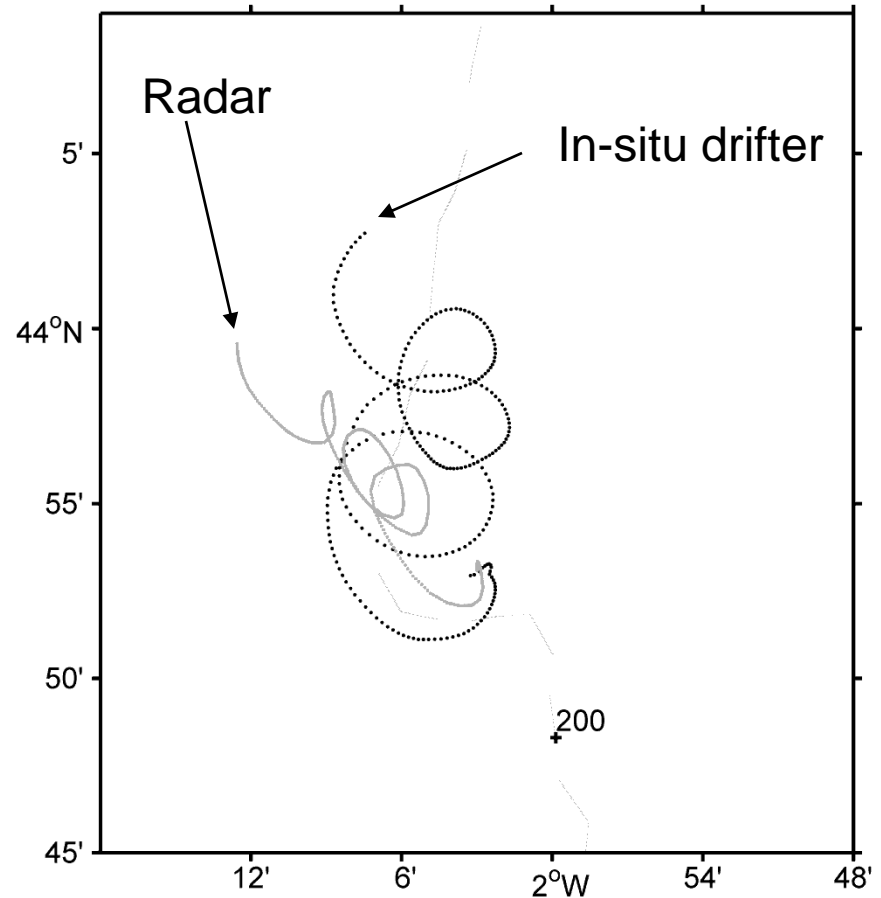
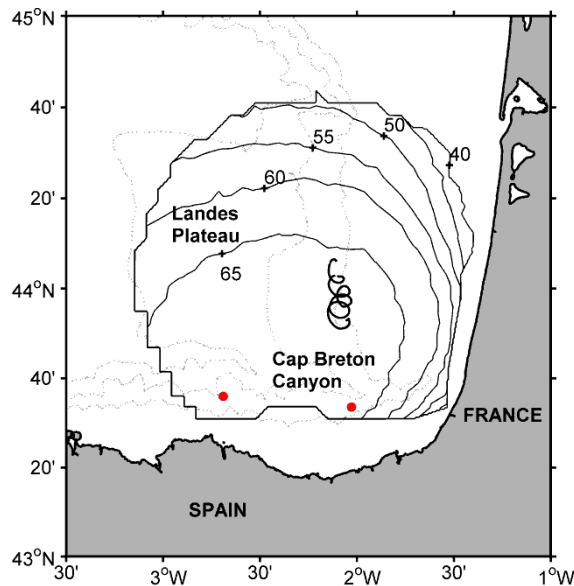
**Radar**



**ADC 1,5 m**

# HF radar vs. In-situ data

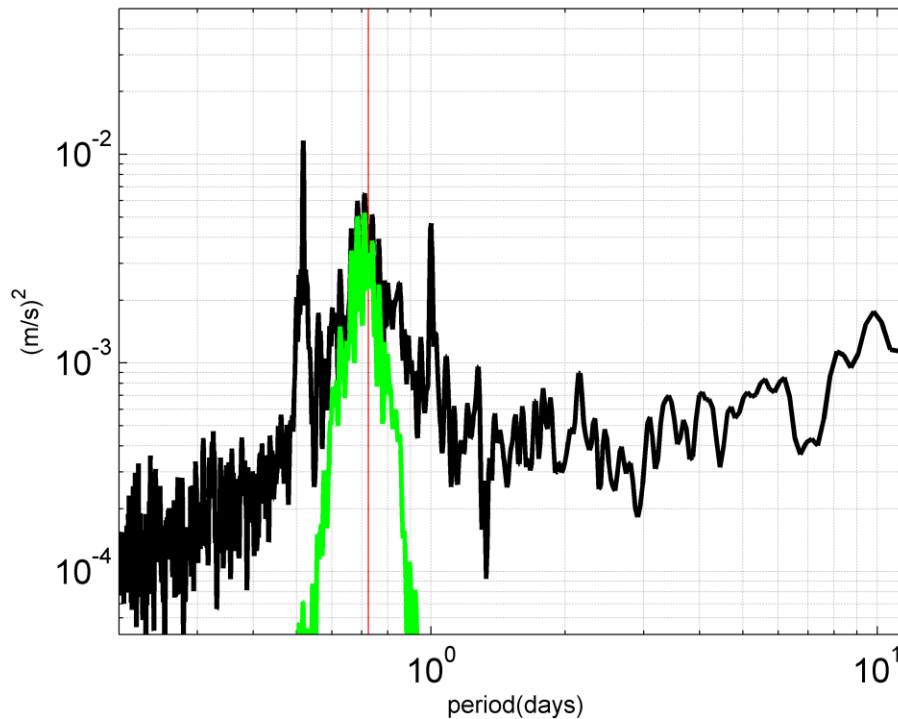
- Dispersion model applied to radar data to reproduce the observed trajectory
- Radar-derived trajectory: smaller amplitude (? Smoothing of data) and mean drift direction rotated to the W (? Ekman)





- 10th order digital Butterworth band pass filter [Emery and Thomson, 2001] applied to each velocity component
  - near-inertial: pass band  $14 \text{ h} < T < 20 \text{ h}$
  - sub-inertial : low pass  $T > 30$

ESPECTRO + FILTROS U 09



Example for the near-inertial pass band

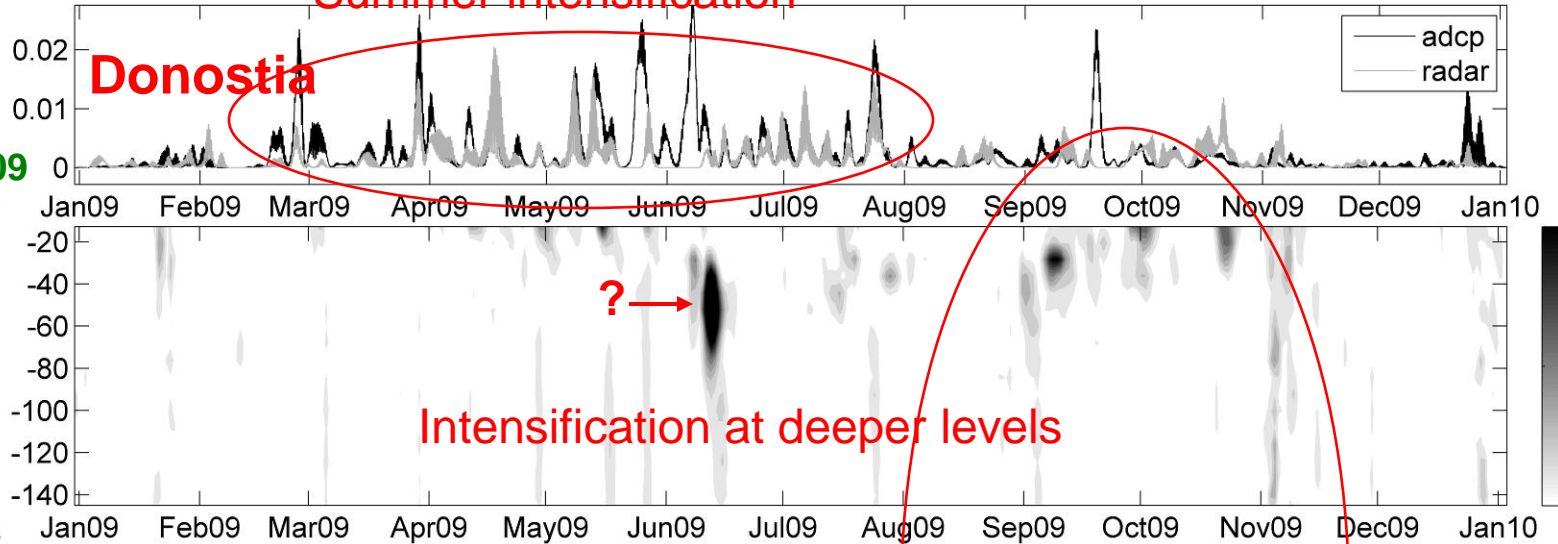
# Distribution of near-inertial KE ( $\text{cm}^2/\text{s}^2$ )

Summer intensification

Surface ADC + RADAR

Donostia

Mean for 2009

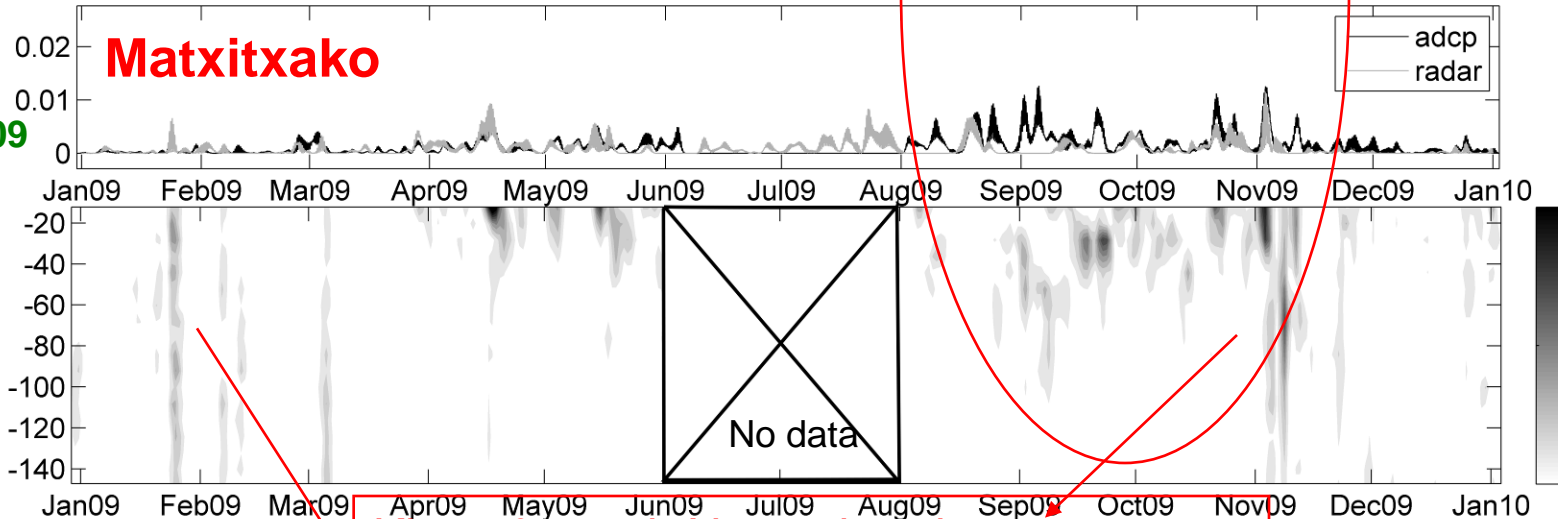


4 6 8 10 12 14  
 $\times 10^{-4}$

Surface ADC + RADAR

Matxitxako

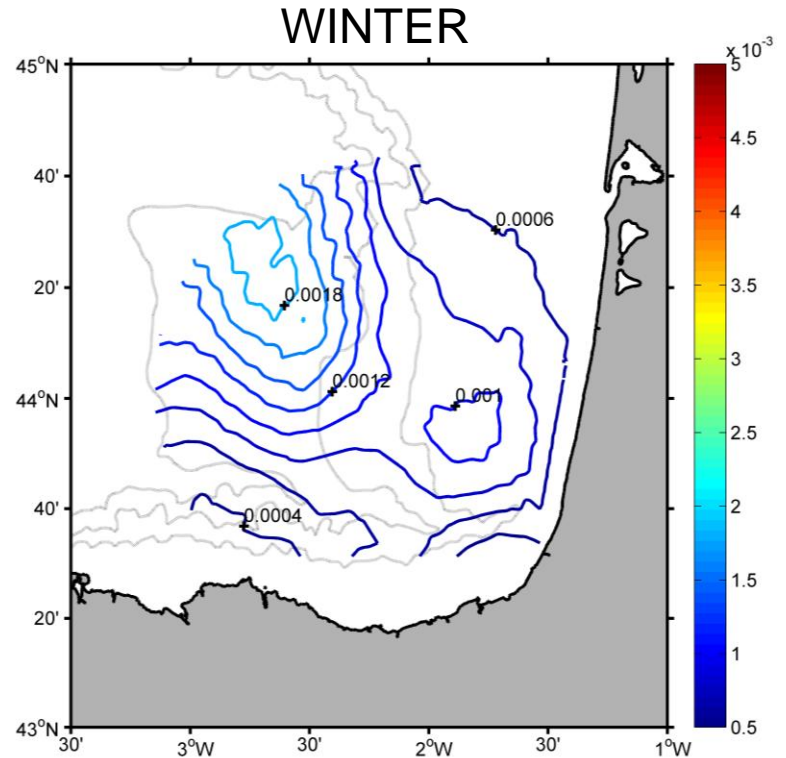
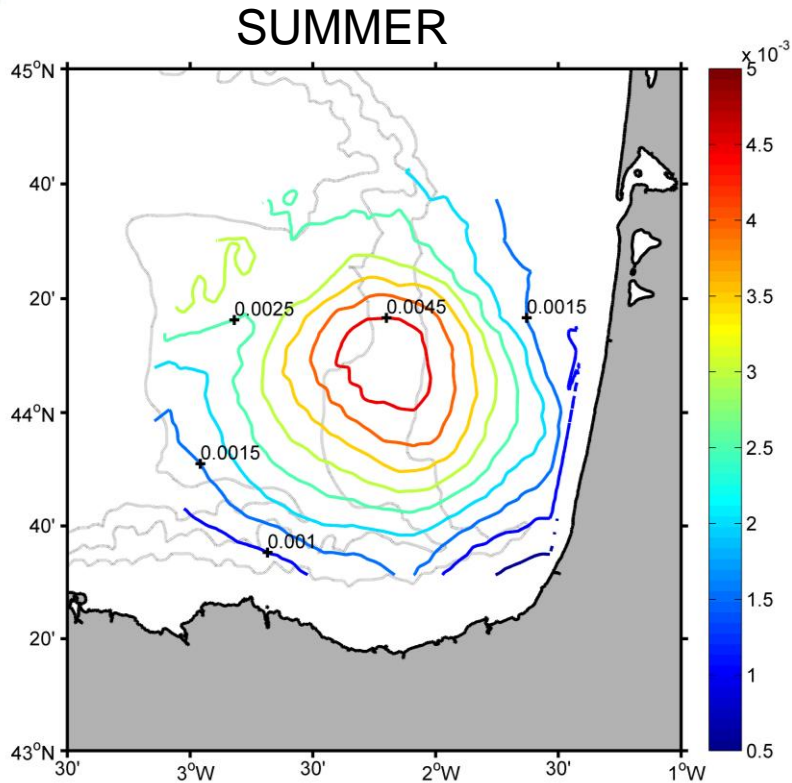
Mean for 2009



2 4 6 8  
 $\times 10^{-4}$

Klaus & 1-10th November intense storms

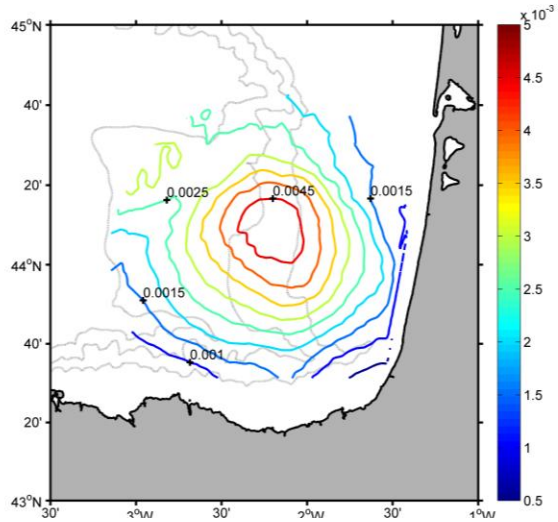
# Distribution of near-inertial KE ( $\text{cm}^2/\text{s}^2$ )



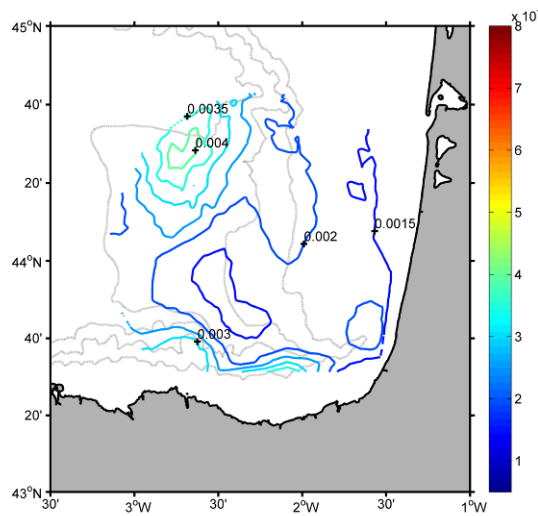
Isobaths: 200, 1000 and 2000 m.

# Distribution of near-inertial KE ( $\text{cm}^2/\text{s}^2$ )- Connection with background field?

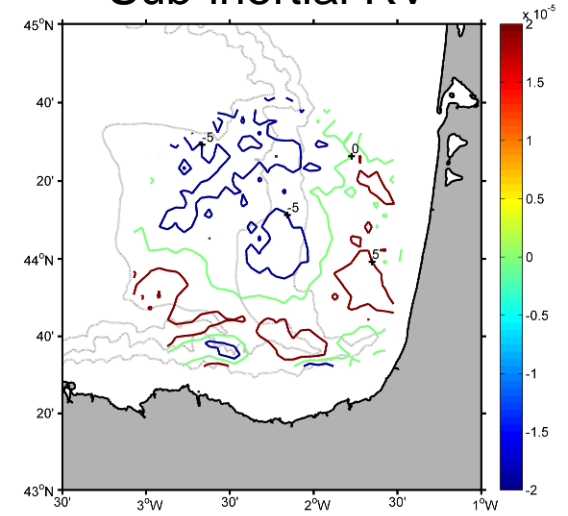
SUMMER



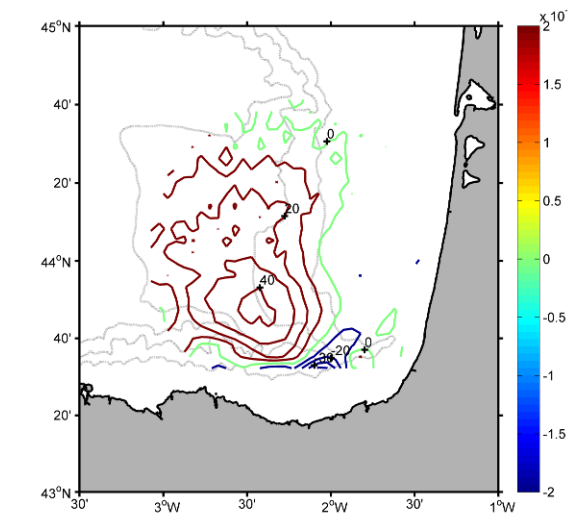
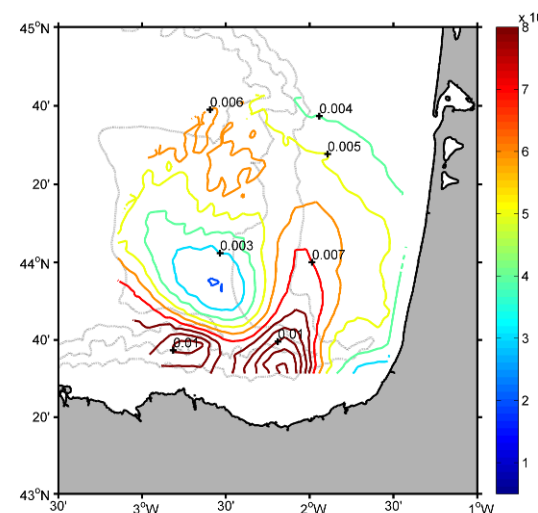
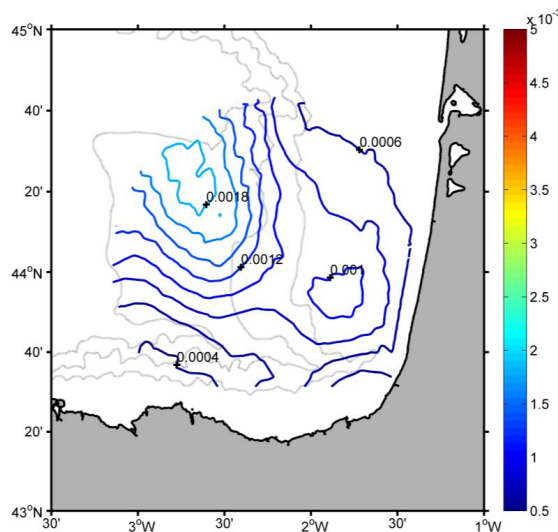
Sub-inertial KE



Sub-inertial RV



WINTER



### SUMMER

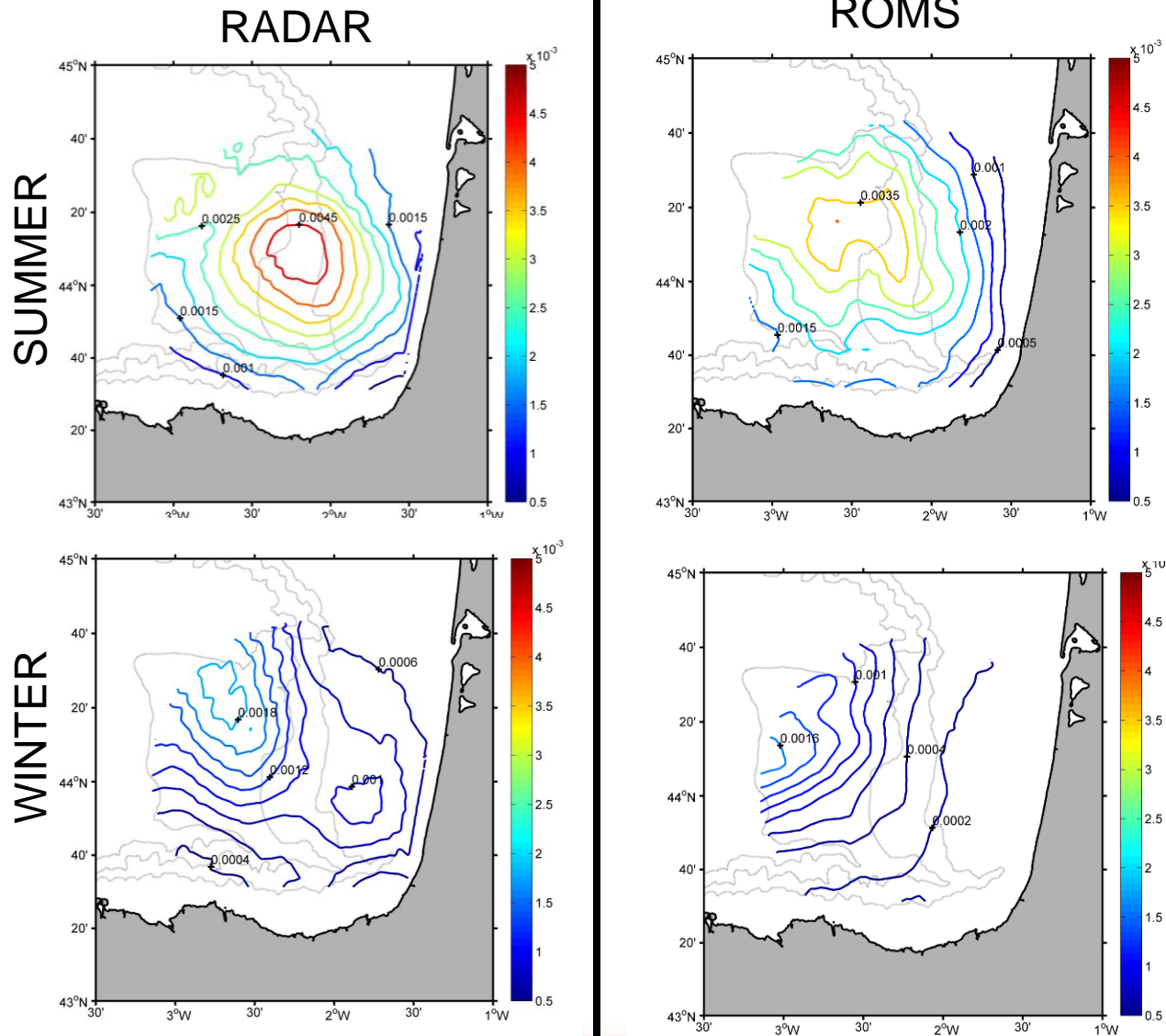
- Weaker and more variable winds, predominantly from the NW
  - anticyclonic pattern with negative vorticity offshore of 200 bathy
  - Stratification, thermocline at 50 m
- The surface near-inertial KE values are much more intense and centred in the middle of the Landes plateau.
  - Surface intensification of the near-inertial oscillations due to stratification.
  - Weak negative RV background field which would favour the oscillations (but it is too weak to explain the observed distribution).

### WINTER

- Intense winds (mainly SW)
  - Well-defined cyclonic pattern, positive vorticity over the slope
  - Well-mixed water column
- Less intense near-inertial KE, confined in the NW of the domain, over the deeper part of the slope
  - Much weaker near-inertial oscillations at surface levels. Deeper mixed layer, and thus less amplitude generation
  - Winter background cyclonic circulation much more intense. The near-inertial oscillations decrease in areas where the mean flow is intense and present strong RV positive values (upper slope).

- ✓ No clear impact of background field (at least in summer when it is weak)
- ✓ Impact of the seasonal differences in atmospheric/buoyancy forcings
- ✓ The summer-winter difference might be also due to wave propagation characteristics ( $C_g$ ) and interaction with coast and bathymetry.

# Next? ROMS/ other model data + other in-situ data (ASPEX)?



**Merci!**

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