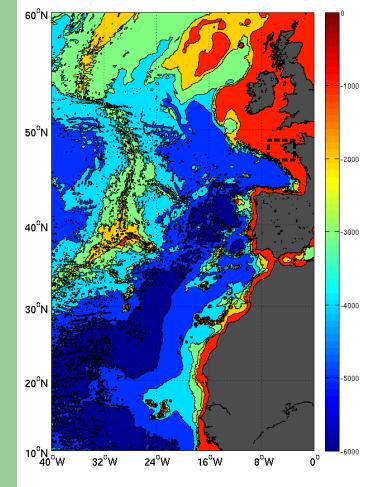
Numerical model and observations of interactions between a continental shelf tidal front and a barotropic jet.

<u>Tanguy SZEKELY</u> Louis MARIE Yves MOREL



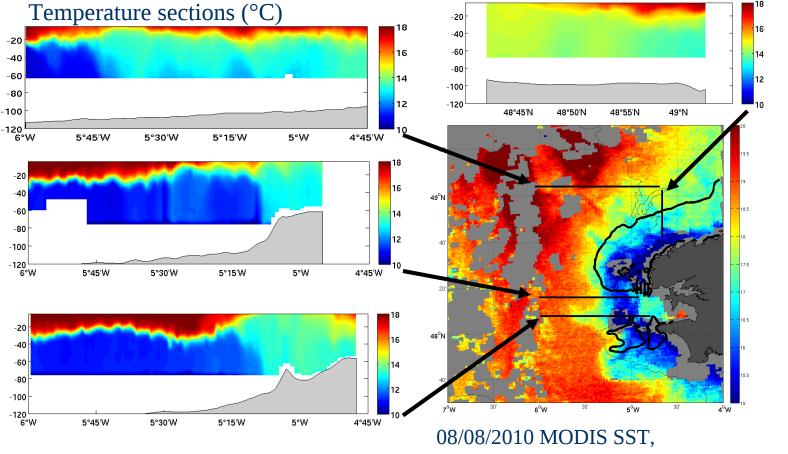
Numerical model and observations of interactions between a continental shelf tidal front and a barotropic jet.



- Introduction, the Iroise sea
- Interactions between a barotropic jet and the baroclinic current induced by a seasonal thermal front (observational).
- Movements of a thermal front driven by the divergence of Ekman layer velocities at the bottom of a jet (numerical).

Hydrography of the Iroise sea: the tidal front

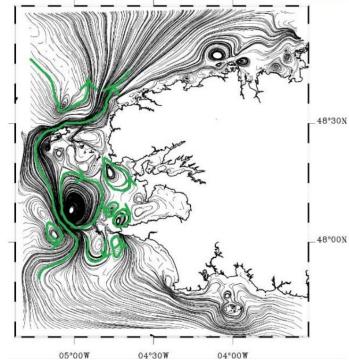
« Scanfish » CTD measurements made during the FROMVAR 2010 cruise.



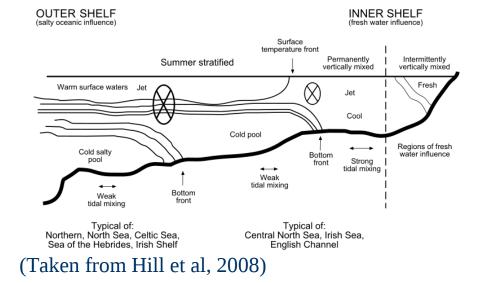
Simpson-Hunter (1974) criterion superimposed

Low frequency circulation

-Semidiurnal **tidal currents** (0.7 m/s) -Strong fortnightly cycle



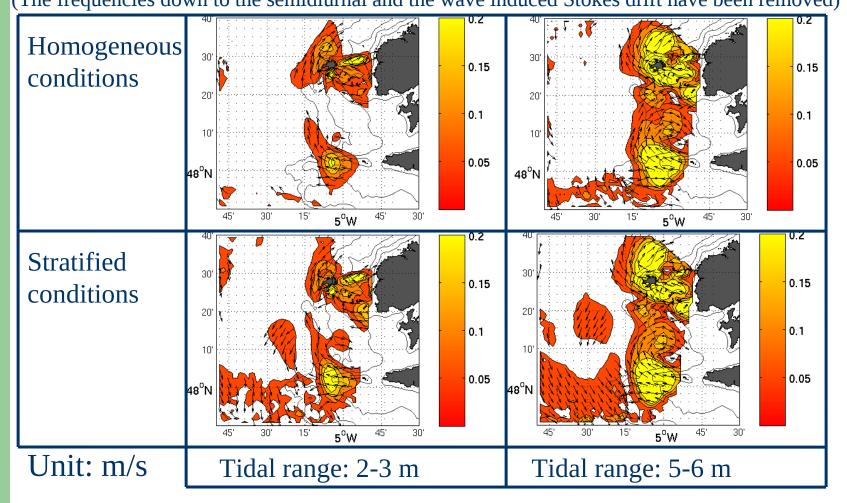
Lagrangian tidal residual streamfunction (M2). Salomon, 1981, Muller 2010



- The northward along-shore low frequency current is composed of both a **tidal residual** and a **thermal wind currents**.

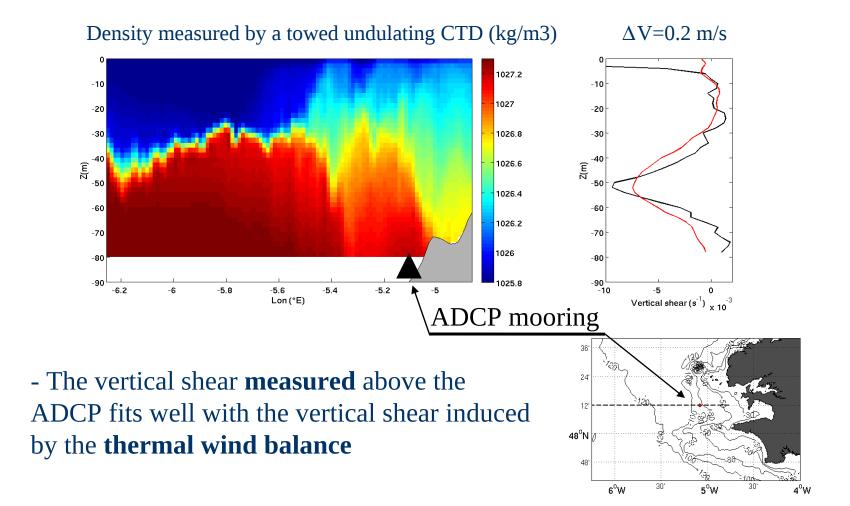
Surface currents measurements

4 years of **surface currents** measurements by two HF radars. (The frequencies down to the semidiurnal and the wave induced Stokes drift have been removed)



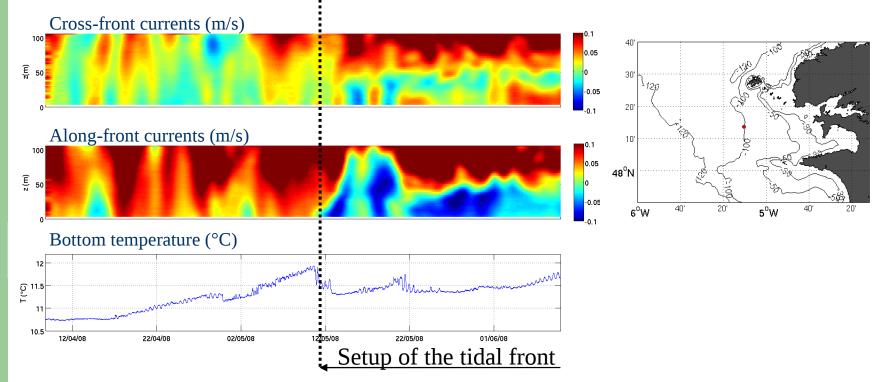
 \Rightarrow Stratification makes no difference in the surface currents. Thermal wind????

Vertical shear measurements



Interactions tides/frontal-current

Tide-filtered velocity measured by an ADCP (April-July 2008)



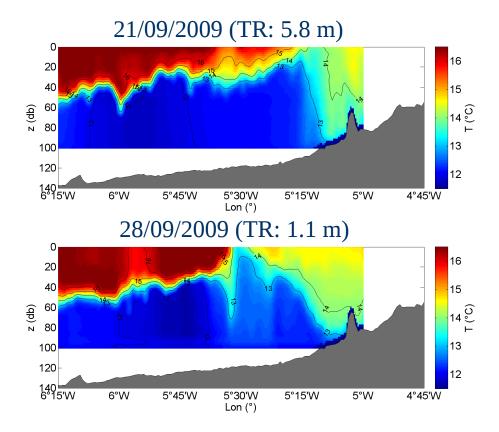
- The low frequency current is **almost barotropic** during winter and **baroclinic** during summer, **but** :
- Thermal wind induced currents are to be sought at the **bottom** !

Part 1: conclusion (observations)

- During **winter** the circulation is dominated by the **barotropic** tidal residual (in periods of low winds).
- During **summer**, the surface circulation remains the same.
- The tidal front induces a **vertical shear** in the along front direction, and a **southward** along front circulation at the **bottom**.

Movements of the bottom front

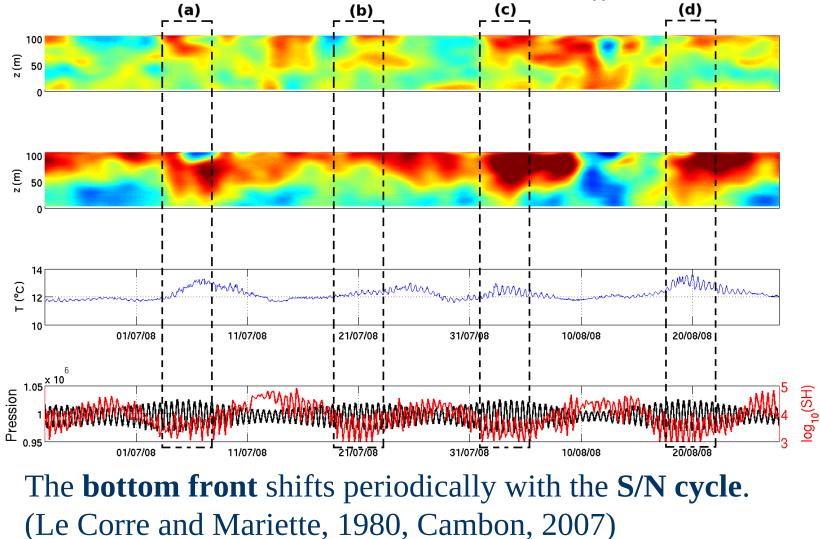
Temperature sections measured by a towed undulationg CTD



The **bottom front** shifts periodically with the **S/N cycle**. (Le Corre and Mariette, 1980, Cambon, 2007)

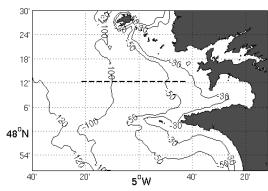
Movements of the bottom front

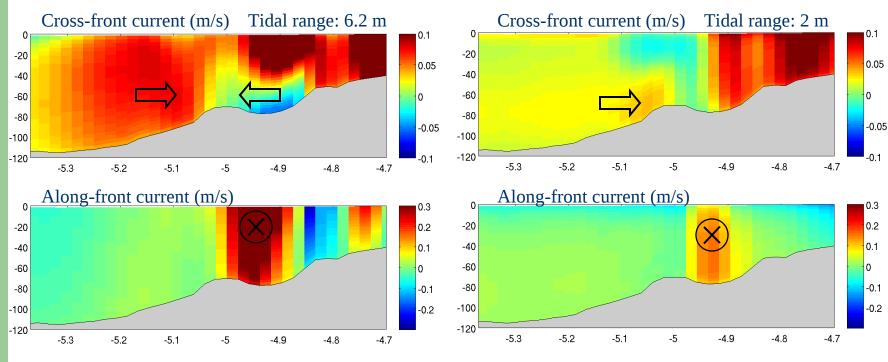
Western and Northern currents from and ADCP mooring.



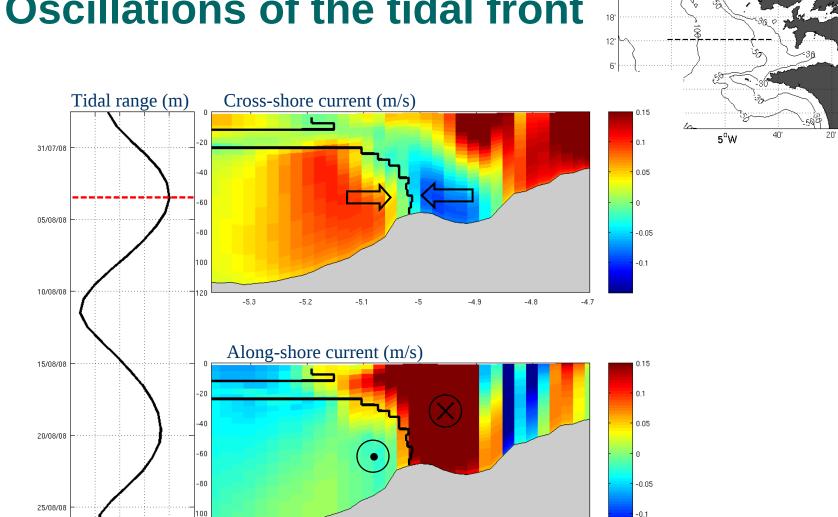
Modelised circulation during homogeneous conditions

ROMS model (Cambon, 2007) : 1.5 km grid, 30 sigma levels Realistic forcing : Meteo France atmospheric/radiative forcings COMAPI tides (F.Lyard, LEGOS)





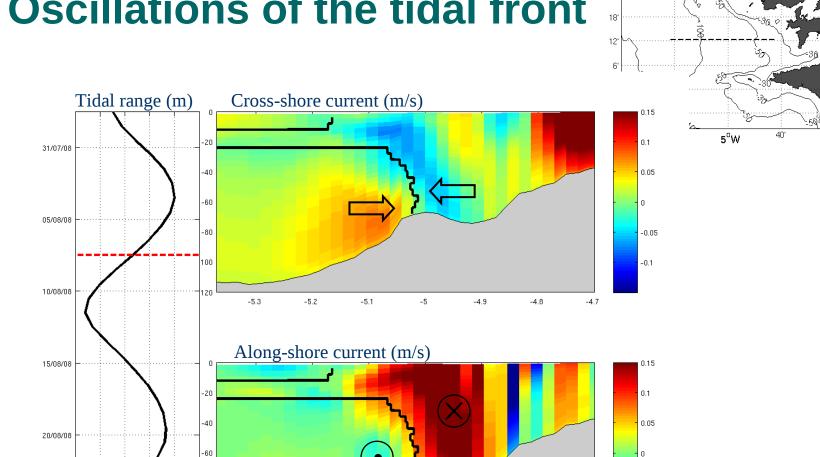
(plots generated without wind forcing, for a homogeneous water column)



30' 24'

Oscillations of the tidal front

120 -5.1 40 60 80 -5.3 -5.2 -5 -4.9 -4.8 20 100 120 -4.7



30' 24'

-0.05

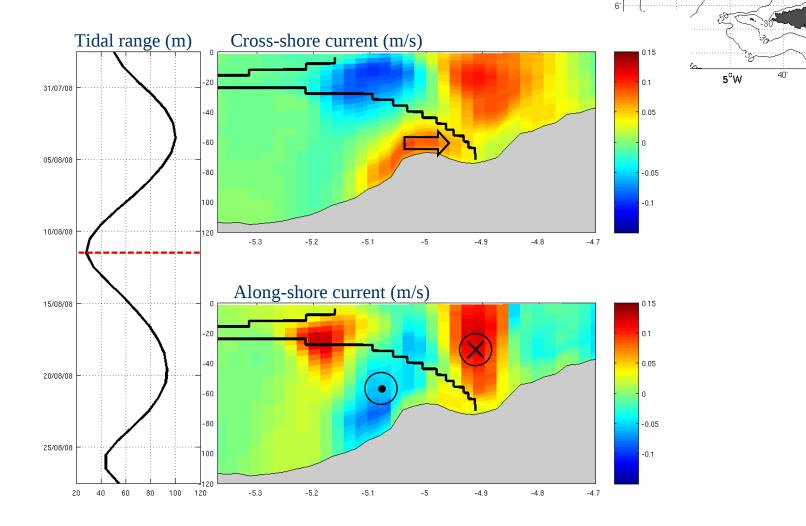
-0.1

20'

Oscillations of the tidal front

-80 25/08/08 100 -5.2 40 60 80 100 -5.3 -5.1 -5 -4.9 -4.8 -4.7 20 120

Oscillations of the tidal front

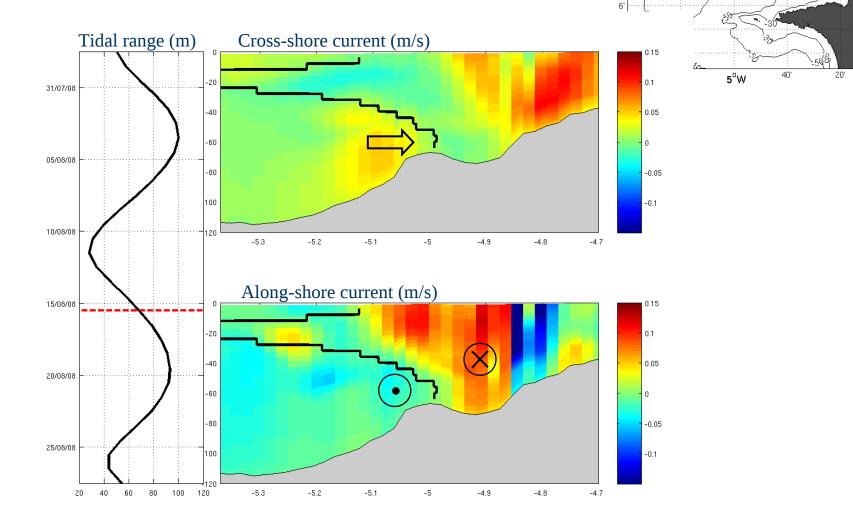


30' 24'

> 18' 12'

> > 20'

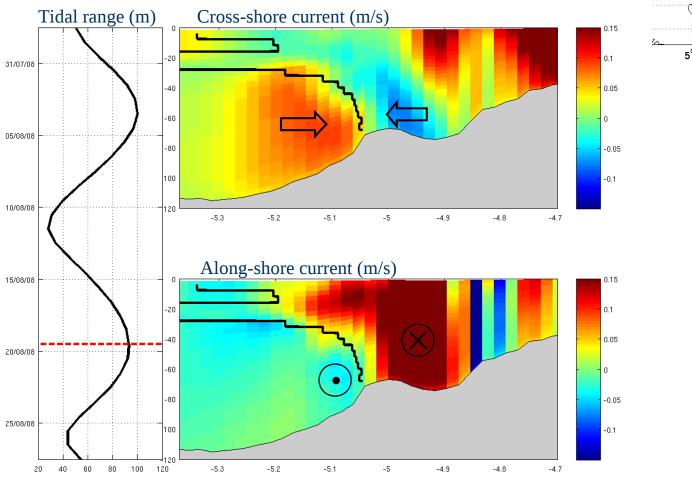


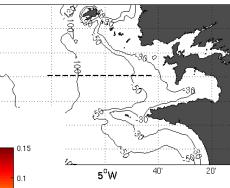


30'

12'

Oscillations of the tidal front





30' 24'

> 18' 12' 6'

Part 2: conclusion (modelling)

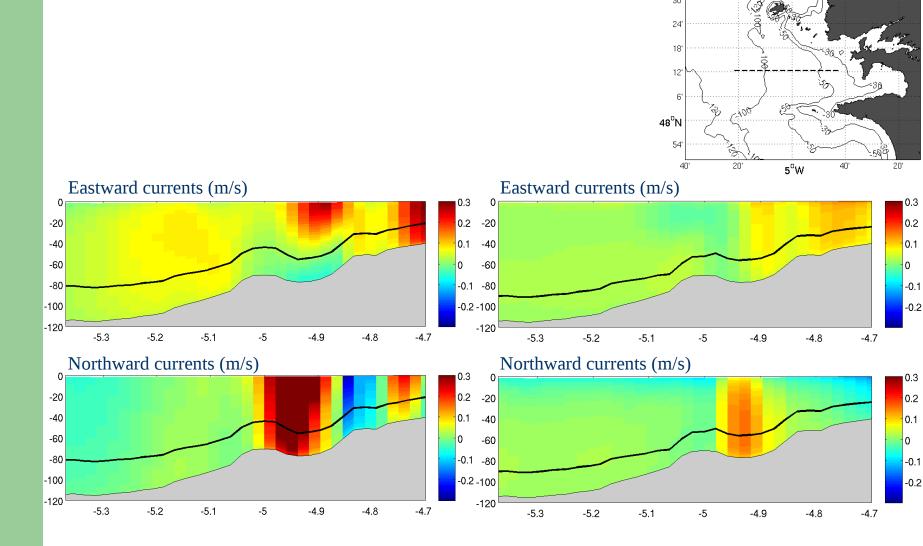
- The variation of the bottom front position is related to the variation of the amplitude of the tidal residual jet.
- Velocities in the Ekman layer generated by the tidal residual jet prevent the collapse of the bottom front.

Conclusions

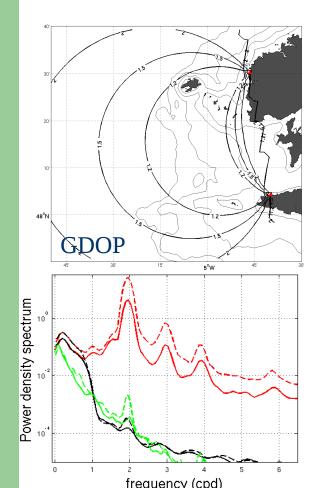
- The bottom front induce **bottom southward** velocities.
- The position of the front is related to the convergence of the velocities in the Ekman layer.

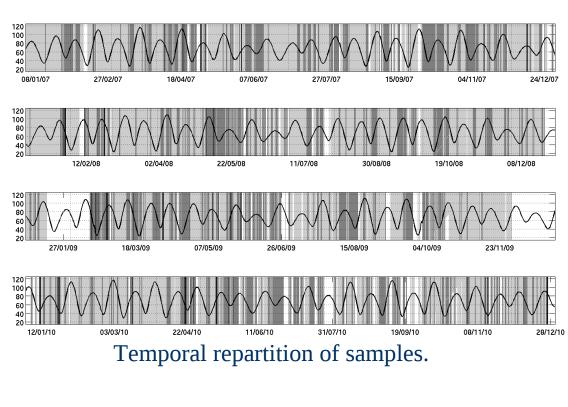


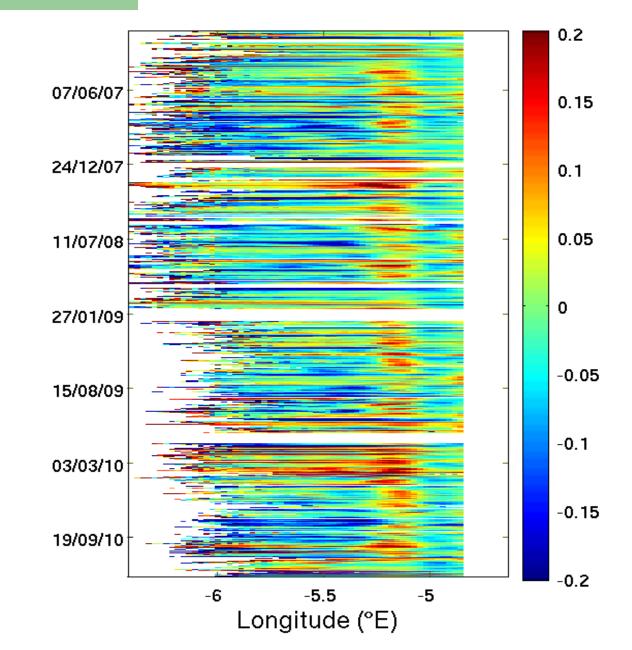
Scaling of the Ekman layer depth



HF radar measurements

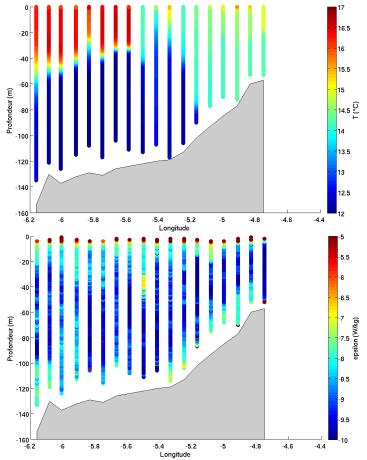


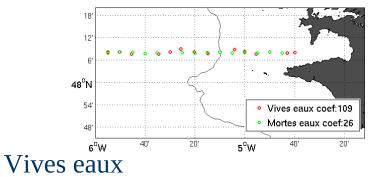


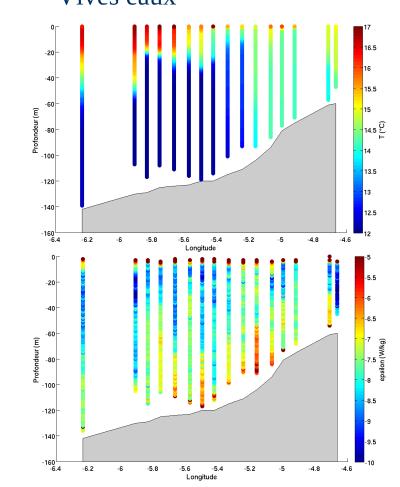


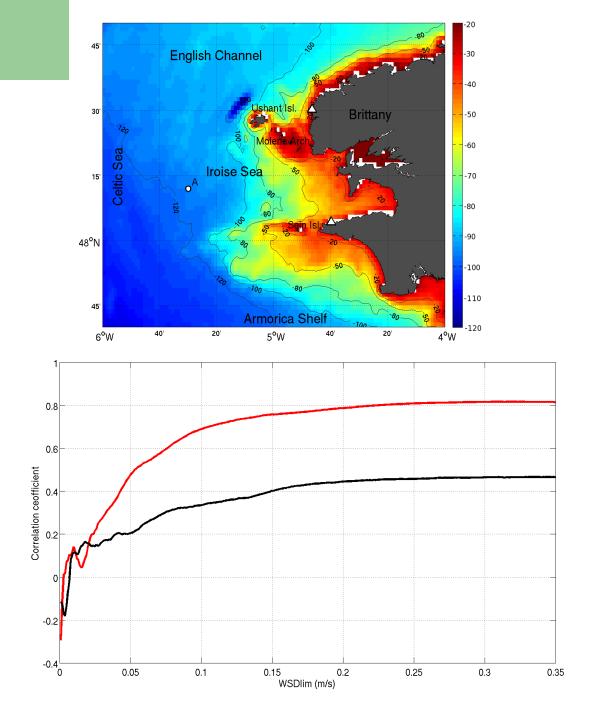
Mesures de microstructure

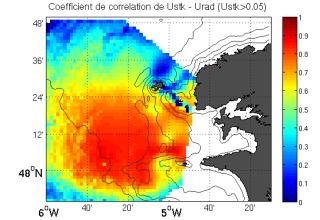




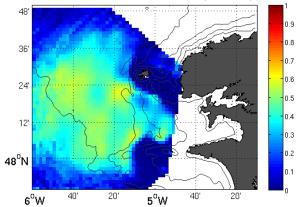








Coefficient de correlation de Ustk - Urad (Ustk<0.05)



Coefficient de correlation de Vstk - Vrad (Ustk>0.05)

