# Bottom friction optimization for barotropic tidal modelling: twin experiments



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# Introduction

- Bottom friction generally represented by a quadratic or a linear formulation : 50°N
  - $\vec{\tau} = -\rho C_D |\vec{u}| \vec{u}$  $\vec{\tau} = -\rho C_L \vec{u}$
- C<sub>D</sub> and C<sub>L</sub> often
   considered as
   constants



 Need for a more physically consistent approach in coastal areas

- HYCOM in a pure barotropic configuration (one isopycnal layer)
- Horizontal resolution: 5.5 km
- Sea-surface elevations and 2D velocities (from the NEA2004 tidal atlas) prescribed along the open boundaries

• Bottom drag formulation : 
$$\vec{\tau} = \frac{-C_D |\vec{u}| \vec{u}}{H}$$

Friction coefficient computation based on vertical integration of the turbulent velocity profile:

$$C_{D} = \left(\frac{\kappa}{\ln\left(\frac{H}{z_{0}}\right) - 1}\right)^{2}$$

κ : Von Kármán constant
 z<sub>o</sub>: bottom roughness
 H : water height

- Simultaneous Perturbation Stochastic Approximation (SPSA, Spall (1998))
- Iterative algorithm :  $\hat{\theta}_{k+1} = \hat{\theta}_k a_k \hat{g}_k (\hat{\theta}_k)$
- Gradient estimation with two loss function measurements :

$$\hat{g}_{ki} = \frac{J(\hat{\theta}_k + c_k \Delta_{ki}) - J(\hat{\theta}_k - c_k \Delta_{ki})}{2c_k \Delta_{ki}}$$

$$\Delta_{ki} = \begin{cases} 1, \text{ probability } 1/2 \\ -1, \text{ probability } 1/2 \end{cases}$$

 $\rightarrow$  Computational cost per iteration independent of the dimension problem

- "Observations" obtained with direct model integration
- Estimated parameter: *z<sub>o</sub>* (bottom roughness)
- Considerable problem dimension: parameter estimation only for some points (hereafter colocation points), and 2D reconstruction using interpolation
- Description of an iteration :
  - 2 perturbed runs and associated loss function evaluations
  - Gradient computation and updating of the estimated parameter
  - 1 run with the new parameter distribution to evaluate the improvement or the deterioration in terms of loss function

- "Observations" obtained with a uniform  $z_o$  distribution :  $z_o = 8mm$
- Only one tidal component: M2 (lunar semi-diurnal, most important component for the area)
- Uniform variation of z<sub>o</sub> (only one colocation point)
- Choices for *a<sub>k</sub>* and *c<sub>k</sub>* :

0 0 0 0

$$c_{k} = \frac{0.005}{k^{1.2} + 10} \qquad \hat{g}_{ki} = \frac{J(\hat{\theta}_{k} + c_{k}\Delta_{ki}) - J(\hat{\theta}_{k} - c_{k}\Delta_{ki})}{2c_{k}\Delta_{ki}}$$

$$a_{k} = \frac{0.003}{k^{0.6} + 10} \qquad \hat{\theta}_{k+1} = \hat{\theta}_{k} - a_{k}\hat{g}_{k}(\hat{\theta}_{k})$$

- Assimilation window : 3.25 days
  - Two days to enable the system to be in equilibrium with the imposed perturbation
    Loss function computation during two periods of M2 (24.84 h)
- Loss function calculated with SSH only (observations are supposed to be available at each grid point)

$$J = 0.5 \sum_{t=1}^{T} \sum_{i=1}^{N} (SSHmod_{i,t} - SSHobs_{i,t})^{2}$$

 Need for a modification of the algorithm: gradient normalization when parameter is updated (gradient is normalized by the maximum gradient of the previous iterations)

#### Two cases:

- start with  $z_0 = 5 \text{ mm}$
- start with  $z_0 = 11 \text{ mm}$



Final value :  $z_0 = 7.999966 \text{ mm}$ Final value :  $z_0 = 8.000005 \text{ mm}$ 



- 42 colocation points
- Start with uniform  $z_0 = 5mm$
- Modified perturbations amplitudes ( $c_k$ ): standard deviation of the parameter for the last N iterations (here N=20)





• "Observations" obtained with a variable  $z_0$  distribution (gaussian distribution)







- SPSA algorithm easy to implement and appropriate for this study
- Realistic experiments:
  - use of diverse data sources: tide gauges, ADCP moorings, satellite data (coastal products), HF radars
  - link between optimal z<sub>0</sub> distribution and a physical reality
  - quantification of the contribution of specific data, especially HF radars in the Iroise Sea